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# Atmospheric modeling and source reconstruction of radioactive ruthenium from an undeclared major release in 2017

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**In October 2017 unusual  $^{106}\text{Ru}$  detections across most of Europe prompted the Institut de Radioprotection et de Sûreté Nucléaire (IRSN) to analyze the event in order to locate the origin and identify the magnitude of the release. This paper presents the inverse modeling techniques used during the event to achieve this goal. The method is based on a variational approach and consists of using air concentration measurements with the Ix long-range dispersion model included in the IRSN's C3X operational platform. The method made it possible to quickly identify the southern Urals as the most likely geographical origin of the release. Despite uncertainties regarding the starting date of the release, calculations show that it potentially began on 23 September, while most of the release was emitted on 26 September. Among the nuclear plants identified in the southern Urals, the Mayak complex is that from which the dispersion of the  $^{106}\text{Ru}$  plume is most consistent with observations. The reconstructed  $^{106}\text{Ru}$  source term from Mayak is  $\sim 250$  TBq. In total, it was found that for 72% of the measurements simulated and observed air concentration agreed within a factor of 5. In addition, the simulated deposition of  $^{106}\text{Ru}$  agrees with the observed deposition. Outside the southern Urals, the simulations indicate that areas with highest deposition values are located in southern Scandinavia and southeastern Bulgaria and are explained by rainfall events occurring while the plume was passing over.**

inverse modeling | ruthenium detection | atmospheric dispersion modeling | source reconstruction

A rare episode of low levels of particulate radioactive ruthenium ( $^{106}\text{Ru}$ ) was observed in the atmosphere of most European countries between late September and mid-October 2017. The concentrations reported were in the range of several microbecquerels per cubic meter to more than  $100 \text{ mBq/m}^3$  (1). As for the  $^{131}\text{I}$  detection event in January/February 2017 (2), no elevation in gamma dose rate was reported by measurement networks around known nuclear sites in Europe. While this makes a major accidental release scenario unlikely, analyzing only activity measurements in the air, for which the observed values are usually weekly averaged, cannot identify either the origin or the magnitude of the release. For such situation, the Institut de Radioprotection et de Sûreté Nucléaire (IRSN) uses modeling techniques to analyze the event in more detail. The aim is, in particular, to identify the origin and to assess the duration and the magnitude of the releases.

The first methods used to estimate the source location of a nuclear release were empirical and consisted of the analysis of back-trajectories and retro-plumes (3). The current methods, which are more efficient, are based on inverse modeling techniques, which combine environmental measurements with atmospheric dispersion modeling and are based on a rigorous mathematical formalism. Inverse methods include Bayesian approaches in which probabilistic considerations may be introduced into the problem in order to account for uncertainties for the input data. Very popular Bayesian techniques are random

search algorithms such as Markov chain Monte Carlo (MCMC) methods (4–6). MCMC methods can be applied to major accident situations such as the Fukushima and Chernobyl accidents. In that context, the aim is to estimate how the releases will evolve with time (7), since the location is known. MCMC methods are also relevant for radionuclide detection events of smaller magnitude where the release location is unknown. As regards the problem of source identification, an MCMC method was applied to the Algeciras incident (8) that occurred at the end of May 1998 and led to a radioactive release following the fusion of a  $^{137}\text{Cs}$  radioactive medical source. In fall 2011,  $^{131}\text{I}$  was detected at several monitoring stations in central Europe following a  $^{131}\text{I}$  release from the Institute of Isotopes in Budapest. The source location, the magnitude, and the temporal evolution of the release are retrieved in ref. 9 assuming no prior information. A Bayesian method was used for source reconstruction for real-world activity concentration data measured by the International Monitoring System radionuclide network maintained under the auspices of the Comprehensive Nuclear-Test-Ban Treaty Organization (10). Variational inverse modeling methods (11–16) are a variant of Bayesian methods since they draw on the Bayes formula but consist only of estimating the optimal solution and not obtaining the probability density function of the estimated source parameters. A variational method was applied to identify the origin of  $^{131}\text{I}$  detected in Europe between January and February 2017 (2). As no information on the source was available, several potential releases

## Significance

**In October 2017, most European networks involved in the monitoring of atmospheric radioactive contamination measured small  $^{106}\text{Ru}$  air concentrations. One of the particularities of the event was that the location and the magnitude of the source were unknown at the time when the first detections were reported. This paper describes the use of inverse modeling techniques combining an atmospheric transport model with environmental measurements in order to identify the origin of the  $^{106}\text{Ru}$  detections and to assess the magnitude of the  $^{106}\text{Ru}$  emissions. The method, easily embedded for operational use, made it possible to quickly point out the southern Urals in the Russia Federation as the most likely geographical origin of the release.**

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The authors declare no competing interest.

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sites were considered to reduce the size of the inverse problem. Compared to variational approaches, MCMC Bayesian methods may be not suitable for operational use because of their prohibitive computational costs, especially where there is a large number of source parameters to sample. Another type of inverse modeling method is maximum entropy on the mean (17–19). It is based on a formalism which is a variant of the Bayes approach, enabling prior information, for example the source positivity, to be taken into account. Maximum entropy of the mean offers a general framework in which the information input prior to inversion is used in a flexible and controlled way. The inversion is shown to be equivalent to the minimization of an optimal cost function, expressed in the dual space of observations. It has been shown that such techniques are efficient in the case of an accidental event like Chernobyl (20).

In this paper, a variational inverse modeling method is used to determine the most probable  $^{106}\text{Ru}$  source location and its magnitude using  $^{106}\text{Ru}$  air concentration measurements. The decision to use a variational inverse technique was based on its ease of implementation and the fact that the computing time is less prohibitive than with Bayesian methods. We first describe the observations reported in Europe. Then, the methodology of the inverse method is presented and applied to identify the  $^{106}\text{Ru}$  source location and the quantities released in the atmosphere from the most reliable location. The main results obtained by inverse modeling are discussed and validated by performing a model-to-data comparison.

## Observations

**Air Concentration Measurements.** The available dataset of  $^{106}\text{Ru}$  air concentrations (for further details see ref. 1) was mainly compiled by Ro5 reports and was completed with data from the Roshydromet website (21), the Typhoon Association website (22), and the EGASMRO website (23, 24). Ruthenium-106 was observed between the end of September and mid-October 2017 in the atmosphere of 31 countries on the European continent at levels ranging from a few microbecquerels per cubic meter to more than 170 mBq/m<sup>3</sup>. Only stations located in western Europe (Portugal, Spain, Great Britain, Benelux, and Northern Ireland) did not report detections of  $^{106}\text{Ru}$  above detection limits. Outside Europe,  $^{106}\text{Ru}$  was also detected at very low levels (microbecquerels per cubic meter) in Guadeloupe, Kuwait, Florida, the eastern part of the Russian Federation, and Mongolia. The detection limit of the measurements also varied significantly between stations depending on airborne concentration, sampling equipment, and measurement capabilities. As outlined in ref. 1, maximum air concentration values were measured in Romania (176 mBq/m<sup>3</sup>), which was explained by the variability of air sampling duration in various countries. The air sampling periods differ from one country to another, ranging from half a day to several weeks. In addition, a given result may correspond to the compilation of several sub-period filters leading to a composite filter sample corresponding up to several months. Hence, it was not possible to identify the source release location solely via the average value over a given sampling period.

**Deposition Measurements.** In addition to the air concentration measurements, deposition was also measured in a number of European countries (see ref. 1). They were either samples taken after the releases were over, giving the cumulative  $^{106}\text{Ru}$  deposit at a particular location, or samples showing the total quantity of  $^{106}\text{Ru}$  deposited in a single day. The daily samples are of greatest interest because they make it possible to work out the period when the  $^{106}\text{Ru}$  was deposited. A number of daily deposition observations were recorded in the Russian Federation, between the Urals and the regions further west, as well as in the rest of Europe.

Although the number of deposition reports is much smaller than the number of air concentration measurements, analysis of the daily deposit observations identifies southern Ural as the first region to have detected the presence of  $^{106}\text{Ru}$ . Daily samples on 23 September in Kyshtym, on 25 September in Argayash, and on 26 September in Bugulma, Dema-Ufa, Metlino, and Novogornyy indicate that  $^{106}\text{Ru}$  deposits reached several hundred becquerels per square meter. Several samples were also taken in December 2017 around the Mayak site in southern Ural. One of the samples situated around 15 km southwest of Mayak confirmed the presence of  $^{106}\text{Ru}$  with deposits of between 500 and 1,200 Bq/m<sup>2</sup> (25). The levels of  $^{106}\text{Ru}$  deposits were much lower at locations west of the Urals (maximum of 17 Bq/m<sup>2</sup> in Morozovsk, 17 Bq/m<sup>2</sup> in Dema-Ufa, and 30 Bq/m<sup>2</sup> in Bugulma). Elsewhere in Europe, the highest deposits were recorded in Scandinavia with levels of up to 90 Bq/m<sup>2</sup> in Finland and 45 Bq/m<sup>2</sup> in Sweden (26). Deposition of  $^{106}\text{Ru}$  was also reported from Poland, Austria, Italy, and the Czech Republic, with a few becquerels per square meter of  $^{106}\text{Ru}$  (1).

## Source Reconstruction Methodology

In this section, we introduce the methodological framework of the source reconstruction used in this study, enabling the geographical origin of the emissions and the  $^{106}\text{Ru}$  quantities emitted in the environment to be determined. The method is based on variational inverse modeling techniques and inspired by that of ref. 27. Although daily deposition measurements provide useful information about the starting date of the release, these are few and are not used during the inversion process. Inverse modeling approaches combining air activity and deposition measurements simultaneously have already been implemented to reconstruct the source term from the Fukushima accident (13, 15). However, these methods require the estimation of additional parameters quantifying the confidence level associated with each type of measurement, which makes their real-time use difficult. Therefore, only air concentration measurements are used in this study. The implementation of the source reconstruction methodology consists of 3 main steps.

### A Priori Information

Because  $^{106}\text{Ru}$  is not normally detected in the environment, it is reasonable to think that one single source of release caused the detections of  $^{106}\text{Ru}$ . The hypothesis of concomitant releases of  $^{106}\text{Ru}$  was therefore not taken into account in this study. Because of apparent west-to-east gradient in  $^{106}\text{Ru}$  detections, it is assumed that the release occurred somewhere between western Europe and the Russian Federation. Dimensions of the domain containing the source are [10W, 70E], [34N, 70 N], covering the majority of the detections of  $^{106}\text{Ru}$  (except those situated outside Europe). For computation time reasons, the initial domain is subdivided into regular mesh cells of resolution  $2^\circ \times 2^\circ$ , leading to a total of 720 mesh cells. Each cell center  $c_k$  is then considered to be a potential source of release. With no prior on the geographical origin of the  $^{106}\text{Ru}$  detections, the probability that the source comes from one particular cell is the same for all cells of the domain.

### Inverse Modeling Method

For each potential source contained in [10W, 70E], [34N, 70 N] related to the cell center  $c_k$ , the source term  $\sigma_k$  is assessed by inverse modeling using the following variational approach.

**Source–Receptor Relationship.** The variational approach used estimates the source term  $\sigma_k$  using field observations and atmospheric dispersion and deposition modeling. The method is described in several publications (see, for instance, ref. 14). It assumes that the measurement vector  $\mu$  in  $\mathbb{R}^d$  can be described as a linear model

with a source–receptor matrix  $H_k$  (11) and unknown source term vector  $\sigma_k$  in  $\mathbb{R}^N$ :

$$\mu = H_k \sigma_k + \epsilon_k. \quad [1]$$

The  $H_k$  in  $\mathbb{R}^{N \times d}$  source–receptor matrix is the Jacobian matrix of the transport model computed in forward mode under the approach proposed in ref. 28. Each column of  $H_k$  represents the dispersion model's response to a unitary release emitted from the cell center  $c_k$ . In view of the period of the  $^{106}\text{Ru}$  detections, we aim to estimate the daily release rates  $\sigma_k$  from 22 September to 13 October. The number of components of  $\sigma_k$  (i.e., release rate per day in becquerels per second) is therefore equal to  $n = 21$  (number of columns of  $H_k$ ). The vector  $\epsilon_k$  in  $\mathbb{R}^d$  represents a combined model–representativity–instrumental error, hereafter called the observation error.

In general, the source–receptor relationship Eq. 1 constitutes an ill-posed inverse problem and its resolution may fail, particularly when the number of observations is limited. A background (or a priori) term is therefore taken into account when resolving the inverse problem in order to incorporate available knowledge of the source term to be estimated  $\sigma_k$  (28). For the present study, 290 air sampling stations have been considered in the inversion procedure, that is, more than 1,100 measurements. Among all of the stations, some of them did not report  $^{106}\text{Ru}$  traces in the atmosphere. Nonetheless, they provide important information and ensure the inverse problem to be solved is better constrained. The number of air concentration measurements taken into account in the inversion procedure is therefore much higher than the number of the source vector components ( $N \ll d$ ), suggesting that the background term can be omitted (28).

**Modeling Errors.** Observation errors  $\epsilon_k$  defined in Eq. 1 are usually assumed to be Gaussian, following a normal distribution:

$$p(\epsilon_k) = \frac{e^{-\frac{1}{2}(\mu - H_k \sigma_k)^T R^{-1} (\mu - H_k \sigma_k)}}{\sqrt{(2\pi)^d |R|}}, \quad [2]$$

where  $R = E[\epsilon_k \epsilon_k^T]$  is the observation error covariance matrix and  $|R|$  its determinant. A strong disadvantage of Gaussian observation errors is that they give more weight to the high concentration values than to the low values, as it is the value of the model-to-measurement difference that is considered in the probability density function (pdf). One possibility for overcoming this difficulty is to choose a log-normal distribution of observation errors (29) with the following pdf:

$$p(\epsilon_k) = \frac{e^{-\frac{1}{2}(\ln(\mu) - \ln(H_k \sigma_k))^T R^{-1} (\ln(\mu) - \ln(H_k \sigma_k))}}{\sqrt{(2\pi)^d |R|}}. \quad [3]$$

Assuming log-normal observation errors, the application of Bayesian inference leads to

$$p(\sigma_k | \mu) = \frac{p(\mu | \sigma_k) p(\sigma_k)}{p(\mu)} = \frac{p(\epsilon_k) p(\sigma_k)}{p(\mu)} \quad [4]$$

$$\propto \exp \left\{ \frac{1}{2} \ln |R| - \frac{1}{2} (\ln(\mu) - \ln(H_k \sigma_k))^T R^{-1} (\ln(\mu) - \ln(H_k \sigma_k)) \right\}. \quad [5]$$

From this inference, to obtain the maximum a posteriori estimate  $p(\sigma_k | \mu)$ , one should maximize the likelihood  $p(\mu | \sigma_k)$ , which is equivalent to maximizing  $\ln p(\mu | \sigma_k)$  and minimizing the following cost function  $J(\sigma_k)$ :

$$J(\sigma_k) = \frac{1}{2} \ln |R| + \frac{1}{2} (\ln(\mu) - \ln(H_k \sigma_k))^T R^{-1} (\ln(\mu) - \ln(H_k \sigma_k)). \quad [6]$$

$J(\sigma_k)$  measures the log differences between the model predictions  $H_k \sigma_k$  and the real measurements  $\mu$ . As described in ref. 7, the main drawback of log-normal observation errors is that they ascribe too much importance to very low concentration values. One way to mitigate the influence of small concentration values is to introduce a threshold  $\theta$  in the cost function  $J(\sigma_k)$  as follows:

$$J(\sigma_k) = \frac{1}{2} \ln |R| + \frac{1}{2} (\ln(\mu + \theta) - \ln(H_k \sigma_k + \theta))^T R^{-1} (\ln(\mu + \theta) - \ln(H_k \sigma_k + \theta)). \quad [7]$$

This tempers the values of  $J(\sigma_k)$  if there are large differences between the observed and simulated concentrations, particularly for low concentrations with values just above the detection limits. In this paper, we have chosen to consider log-normal observation errors whereas simple parameterization for  $R$  matrix is used. It is assumed that  $R$  is diagonal and the error variance is the same for all diagonal elements (homoscedasticity property):

$$R = r^2 I_d, r > 0. \quad [8]$$

Indeed, minimizing  $J(\sigma_k)$  is equivalent to minimizing

$$J(\sigma_k) = \sum_{i=1}^d (\ln(\mu_i + \theta) - \ln(H_k \sigma_k + \theta)_i)^2. \quad [9]$$

However, the choice of homoscedasticity assumption is not the most appropriate since, theoretically, one should include correlations induced by observation error. More sophisticated approaches are possible (29). The impact of the homoscedasticity assumption will be discussed later in more detail. Furthermore, as in ref. 30, the parameter  $\theta$  is manually selected large enough to prevent the lowest air concentration values from dominating the inverse problem (7, 30). Several values of  $\theta$  were tested:  $\theta = 0.1$ ,  $\theta = 0.01$ , and  $\theta = 0.001$  mBq/m<sup>3</sup>.  $J(\sigma_k)$  is minimized using the L-BFGS-B limited-memory quasi-Newton minimizer (31). The positivity of the source is enforced in L-BFGS-B and no upper bound is used.

### Statistical Indicators

For each potential release at the cell center  $c_k$ , the simulated concentrations at each station  $i$  are trivially given by the product  $(H_k \sigma_k)_i$ . Then, the agreement between simulated and observed air concentration measurements is assessed using the following statistical indicators:

- The reduction factor of the cost function (RFJ) per cell center: Since the initial contribution of each grid center to the cost function has the same value, the RFJ obtained after minimization provides a good indication of the most relevant cells centers for the source location.
- The percentage within a factor of 2 (fac2) per cell center: This is the proportion of the simulated activity concentrations calculated using the reconstructed source vector  $\sigma_k$  that are within a factor of 2 of the observed values. The fac2 indicator is complementary to RFJ since all of the concentration values have the same weight.

The probability that a cell center  $c_k$  is the source location increases with the value of the statistical indicator and therefore the level of agreement between the simulated and observed concentrations. The individual performances of each potential source

are projected by linear interpolation on a map enabling the relevance of the potential release areas to be viewed.

### Sensitivity Analysis

**Sensitivity to Observation Error.** The inverse method based on the likelihood maximization only provides an optimal solution without quantifying the associated uncertainties. One way to quantify this uncertainty is to use a Monte Carlo sensitivity analysis (14). To compute a sample  $n$ , the observation vector  $\mu$  is perturbed such as  $\tilde{\mu}_i = r\mu_i$ ,  $1 \leq i \leq d$ , and  $r$  follows a log-normal law  $r \sim \exp(N(0, \lambda^2))$ , where  $\lambda$  is the SD of the related Gaussian law  $N(0, \lambda^2)$ . Then, for each sample  $n$ , an optimal solution  $\tilde{\sigma}_k$  is obtained by performing the minimization of the cost function  $\tilde{J}(\sigma_k)$ .

**Sensitivity to the Spatial Distribution of Stations.** Assuming that the  $R$  matrix is diagonal suggests that the measurements are not correlated in time and space. In order to quantify the impact of this hypothesis and to evaluate the robustness of the reconstructed source, inverse modeling was applied using  $n$  random subsets of stations with different sizes. The random process used ensures that each station selected is sufficiently far from the others to mitigate the effects of spatial and temporal correlations. Measurements can be therefore considered independent, which is consistent with the assumption of homoscedasticity.

To interpret the results related to sensitivity analysis, the maps of average and SD of the indicators previously defined are plotted.

### Application to $^{106}\text{Ru}$ Source Reconstruction

The aim is to apply the inverse method previously described in order to locate the most probable origin of the  $^{106}\text{Ru}$  detections and the quantities released into the atmosphere between 22 September and 13 October.

### Atmospheric Dispersion Modeling

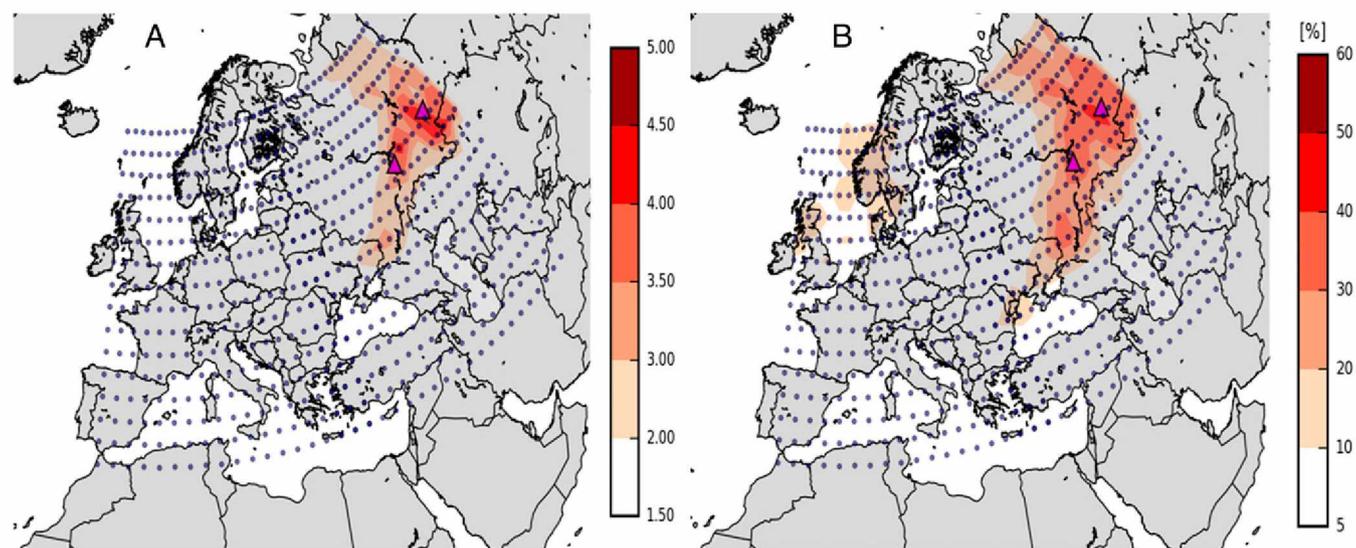
Seven hundred twenty source–receptor matrixes  $H_k$  are computed in accordance with the domain [10W, 70E], [34N, 70 N] of  $2^\circ \times 2^\circ$  resolution. In view of the number of columns of  $H_k$ , it requires launching  $21 \times 720$  forward simulations using the Eulerian atmospheric dispersion model IdX. This model is part of IRSN's C3X operational platform (32). It is based on the Polair3D chemistry transport model (33) and has been validated in

past nuclear accidents (16, 34, 35). IdX takes into account dry and wet deposition as well as radioactive decay and filiation. Dry deposition is modeled by a simple scheme with a constant deposition velocity:  $v_{\text{dep}} = 2.10^{-3}$  m/s. For wet scavenging, the parameterization used is in the form  $\Lambda_s = \Lambda_0 p_0$ , where  $\Lambda_0 = 5.10^{-5}$  h/(mm·s) and  $p_0$  is the rainfall intensity in millimeters per hour (36). The simulations are carried out by forcing IdX with three-hourly operational meteorological data from the ARPEGE model provided by Météo France. The spatial resolution of these data is  $0.5^\circ \times 0.5^\circ$ . The spatial domain of the IdX simulations is [20W, 90E], [20N, 80N], thus encompassing most of the detections. The time resolution is 10 min and the IdX model provides hourly instantaneous concentrations. The spatial resolution is that of the meteorological data ( $0.5^\circ \times 0.5^\circ$ ) and 11 vertical levels between 0 and 4,400 m are considered. The release height is taken to be the first level of the model, that is, between 0 and 40 m.

### Discussion of the Results

**Location of the  $^{106}\text{Ru}$  Release.** The inverse modeling results show that the values of the  $\theta$  parameter have a small influence upon the assessed release location (*SI Appendix, Fig. S1*). For the sake of simplicity, we have chosen to present the results with  $\theta = 0.1$ . Fig. 1 shows a gridded map of the fac2 and the RFJ attached to each potential source. The higher fac2 is, the higher the agreement between modeled and observed air concentrations. The fac2 values are above 40% in a small area in the Russian Federation along the southern Ural Mountains, consistent with ref. 37. Further west, between Ukraine and Volga, the fac2 values are lower and range from 30 to 40%. In addition, analysis of the simulations, considering a source situated in this area, indicates that it is not possible to reproduce the observations in southern Ural between 26 September and 1 October, when several tens of millibecquerels per cubic meter were observed at the Argayash and Novogornyy stations. The winds were blowing from Ural toward Europe throughout this period, so a release situated between Ukraine and Volga, or further west, could not have been carried to Ural. Similarly, a release from Ukraine or further west is not consistent with the detections in Siberia (12).

Even though the inverse modeling results show that the northern Ural region is also a potential source region (fac2 > 30%), the number of air concentration measurements available in this area is



**Fig. 1.** (A) Reduction factor of the cost function. (B) Percentage of simulated air concentrations that are within a factor of 2 of the observed values. Inversion procedure is performed using  $\theta = 0.1$ . Purple triangles represent the location of the Mayak and Dimitrograd sites. Blue dots are the 720 potential source locations.

not sufficient to confirm the relevance of this result. Moreover, the Beloyarsk station situated 150 km north of the Mayak nuclear complex (southern Ural area) did not report the presence of  $^{106}\text{Ru}$  between 18 and 25 September, which is not consistent with a release from the northern Ural region. Therefore, the hypothesis of a release emitted from the southern Ural region remains at this step the most relevant because it better explains all of the measurements. Similar results are obtained when a domain of  $0.5^\circ \times 0.5^\circ$  resolution is considered (*SI Appendix, Fig. S2*).

Beyond this geographical area, the fac2 values decrease rapidly down to below 30% at the Russia–Ukraine border. We also notice that the fac2 values range from only 5 to 20% in Romania, even though the highest air concentration activities were measured in the country. Even further west, the fac2 values fall below 5% nearing Poland, Germany, Spain, and France. This means that a release from western Europe is an unlikely hypothesis. The calculations performed using RFJ indicators restrict the most reliable area of the potential release to the southern Ural region (RFJ > 4.5). In this area, the reduction factor is close to 5, whereas it is close to 2 in the Russia–Ukraine border area. It confirms the results obtained using the fac2 indicator.

The Monte Carlo analysis performed for  $n = 1,000$  samples using realistic value of the SD  $\lambda = 0.5$  indicates that perturbations in observations have a negligible impact on the source reconstruction. Indeed, the highest averaged RFJ values from the 1,000 samples are located in the southern Urals (*SI Appendix, Fig. S3*). In addition, the SD calculated from the RFJ indicator (*SI Appendix, Fig. S3*) is particularly low in the southern Urals, which highlights the robustness of the result.

To assess the effect of the temporal and spatial correlation between stations, the method previously described is applied by considering  $n = 1,000$  random subsets of respectively 10, 20, 50, and 100 stations. A random subset is constructed imposing a minimal distance of 200 km between stations. Then, we applied inverse modeling on each of the random subsets. Since the initial value of the cost function may vary significantly depending on the measurements including in a subset, the use of the RFJ indicator is not suitable to interpret the results. Only the mean and SD of fac2 for subsets of 10, 20, 50, and 100 stations are therefore plotted in *SI Appendix, Figs. S4 and S5*. *SI Appendix, Figs. S4 and S5* show that the southern Urals is identified as the most relevant release area for all subsets considered, including those containing a small number of stations. The higher the size of subset is, the lower the SD of fac2. For example, the SD of fac2 is lower than 5% for a subset of 100 stations, which highlights the robustness of the source location assessment.

**Candidate Sites.** In the area where the fac2 values exceed 30%, to our knowledge 2 nuclear facilities have the capacity to produce sufficiently large quantities of  $^{106}\text{Ru}$  to be detected thousands of kilometers away. They are the Research Institute of Atomic Reactor (RIAR) in Dimitrovgrad (Volga region, 800 km east of Moscow), a factory producing isotopes for medical use, and the Mayak Production Association in Ozyorsk (south of the Urals, 2,000 km east of Moscow), a spent fuel reprocessing facility.

The Mayak site is in the area where the fac2 and cost function reduction factor values are highest (Fig. 1 and *SI Appendix, Fig. S2*). The Dimitrovgrad site is located on the western edge of the most reliable area where the values of fac2 and the factor reduction of the cost function do not exceed 35% and 3.5, respectively. These lower scores are explained partly by the fact that a simulated release from the area of the Dimitrovgrad site does not reproduce the air concentrations in southern Ural measured between 26 September and 1 October. Moreover, the release reconstructed by inverse modeling from the area of the Dimitrovgrad site occurs mainly on 27 September, which is not compatible with the daily deposit observations at Kychtym indicating the presence of  $^{106}\text{Ru}$  from 23 September.

**Source Term Assessment.** For the most relevant area of release defined by fac2 > 30%, the quantity of  $^{106}\text{Ru}$  released estimated by inverse modeling ranges from 100 to 1,000 TBq (*SI Appendix, Fig. S6*). In this area, the release would have started between 23 September in the northern Ural and to a lesser extent 28 September close to the Ukrainian border. In the southern Ural area, identified as the most reliable source location taking account of all of the calculation results, the quantity of  $^{106}\text{Ru}$  ranges from 200 to 600 TBq, consistent with the estimate in refs. 38 and 39. Table 1 describes the estimated quantities from the Mayak site as a function of values of the  $\theta$  parameter.

We notice that the source term varies moderately depending on the values of  $\theta$  and the maximum estimated quantities are 260 TBq for  $\theta = 1$ . In addition, the value of  $\theta = 10^{-1}$  leads to the highest fac2 score (43%) and the maximum error reduction of the cost function (around 78%). The value of  $\theta = 10^{-1}$  is a compromise which avoids giving too much weight to very low concentration values, close to the detection limits, but also avoids giving too much weight to the highest concentrations. In the rest of the paper, we will therefore consider the inversions obtained with  $\theta = 10^{-1}$  as the reference results. In this case, the source term from the Mayak site is  $252 \pm 13$  TBq. Fig. 2 highlights that most of the release occurred on 26 September and very-low-level releases also occurred on 23 and 24 September. Monte Carlo analysis ( $n = 1,000$  inversions using  $\lambda = 0.5$  are performed) ascertains that the quantities estimated between 23 and 25 September do not exceed 2 TBq, whereas the reconstructed quantities released on 26 September are around 250 TBq.

### Model-to-Data Comparison

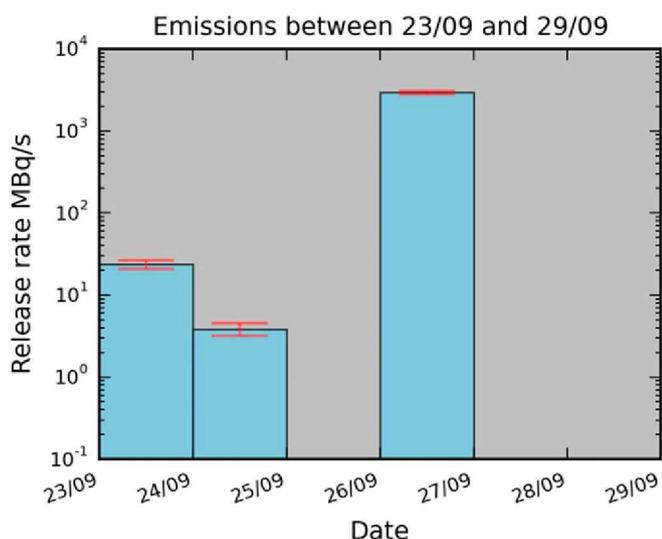
In this section, the source term from the Mayak Production Association is validated by comparing the environmental measurements with the atmospheric dispersion simulation obtained using the IdX model. The agreement between the simulations and the observations over time, the geographical areas, and the type of statistics is evaluated and analyzed.

**Plume Dispersion Analysis.** The simulation of the plume dispersion from the Mayak site is shown in Fig. 3 at different dates between 28 September and 7 October. At the start of this period, the meteorological situation was characterized by the presence of a huge anticyclone centered on northeastern Scandinavia (*SI Appendix, Fig. S7*), encouraging the establishment of an easterly wind between the Urals and a large part of Europe. Following the small releases on 23 September, the first plume reached southeast Europe between 25 and 28 September. The simulated concentration levels in this plume rarely exceed  $1 \text{ mBq/m}^3$  throughout the period. A second plume is then emitted as a result of the large releases on 26 September. Between 26 and 28 September, again driven by the east winds, this plume headed for the Republic of Tartastan in the Russian Federation, which is consistent with the daily deposit measurements in Bugulma indicating the presence of  $^{106}\text{Ru}$  on 26 and 27 September.

No rain is modeled in this geographical area because of the anticyclonic conditions in the region at the time. The plume then

**Table 1. Evolution of the source term, fac2, and the error diminution obtained after minimizing the cost function J depending on the value of  $\theta$  and assuming that Mayak is the release location**

$\theta$ , $\text{mBq/m}^3$	Error reduction of J, %	fac2, %	Source term, TBq
$10^{-3}$	75.2	24	$114 \pm 11$
$10^{-2}$	76.8	37	$189 \pm 15$
$10^{-1}$	77.9	43	$252 \pm 13$
1	75.8	40	$260 \pm 16$



**Fig. 2.** Mean and SD range of the reconstructed release rates (mega-bequerels per second) from Mayak computed using Monte Carlo analysis ( $n = 1,000$  samples). Inversion procedure is performed using  $\theta = 0.1$ .

reached the Rostov Oblast in southwest of the Russian Federation on 28 September, when the deposits of  $^{106}\text{Ru}$  measured at Morozovsk were about  $20 \text{ Bq/m}^2$ , then Ukraine. The plume was over Romania on 29 September. In the middle of the day on 29 September, the highest simulated hourly concentrations exceeded  $100 \text{ mBq/m}^3$  between Ukraine and Romania. During the day on 30 September, the simulated plume split in two: One part spread toward Central Europe and the other moved eastwards toward Siberia. Significant precipitation is modeled in southeast Bulgaria, northern Greece, and the westernmost part of Turkey when the simulated air concentrations become significant ( $>40 \text{ mBq/m}^3$ ). Daily readings from several Bulgarian and Turkish meteorological stations confirm the occurrence of precipitation on 30 September, which agrees with the ARPEGE meteorological model. On 1 October, the main plume was over central Europe and southeastern Europe (Turkey and Greece). The simulated hourly concentrations exceed  $40 \text{ mBq/m}^3$  in the Czech Republic, Serbia, and Bulgaria but were generally lower in Greece and Turkey (between 1 and  $10 \text{ mBq/m}^3$ ). Hourly simulated concentration levels also exceed  $40 \text{ mBq/m}^3$  in northern Kazakhstan, although no data are available to confirm the simulated concentration levels. On 2 October, the main part of the plume went northward toward Scandinavia and reached the eastern part of Italy, where the simulated levels were up to  $40 \text{ mBq/m}^3$ . Because the anticyclone over Scandinavia subsided (*SI Appendix, Fig. S7*), the plume's progress westward was halted. The anticyclone was replaced by a more oceanic weather pattern reflected in the presence of a large rainy front stretching from Germany to Norway at a time when the simulated hourly concentrations were falling but remained close to  $30 \text{ mBq/m}^3$  in southern Sweden. Between 2 and 3 October, the eastern part of the plume continued to move eastward toward Siberia. The simulated concentrations are much more significant ( $>10 \text{ mBq/m}^3$ ) further west toward Kazakhstan. From 3 October, westerly winds settled in over a large area of western Europe after the rainy front had passed, causing the plume to spread eastward and have an impact again on several stations in eastern Europe. Levels gradually fell in Europe as the plume spread to North Africa (Libya and Egypt) and the Middle East. Finally, small precipitations are reproduced by the ARPEGE models in the western part of the Russian Federation on 6 and 7 October, when the simulated concentrations are below  $1 \text{ mBq/m}^3$ .

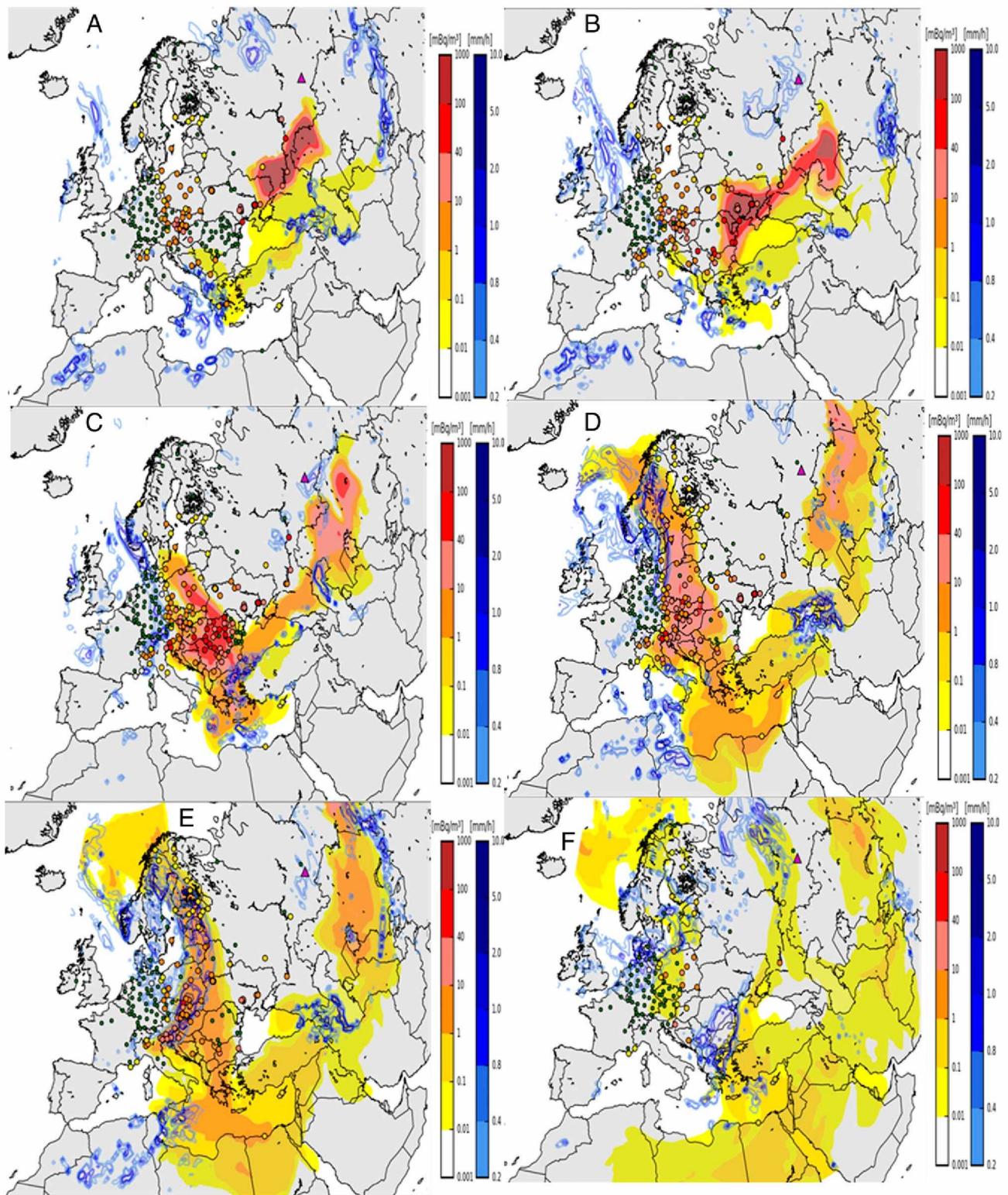
**Comparison with  $^{106}\text{Ru}$  Air Concentration Measurements.** The objective is to study the ability of the simulations to reproduce the observed concentrations and therefore the realism of the assessed source terms. Fig. 4 shows the maximum concentrations observed and simulated by station using the reconstructed source term from the Mayak complex. According to Fig. 4, the simulations clearly distinguish the area delimiting, on the one hand, the stations that did not detect  $^{106}\text{Ru}$ , in the western part of Europe and, on the other the stations further east that measured higher levels of  $^{106}\text{Ru}$ . This area extends from southeastern France to Denmark, passing through Germany. The maximum observed concentrations are generally well reproduced by the simulations, as shown by the fac2 values presented in Table 1. Fig. 5, which provides examples of comparisons between simulated and observed concentrations at different stations, confirms that the simulation performed using the reconstructed source term satisfactorily reproduces all of the observations. The scores obtained are, for example, higher than for other accident situations such as the Fukushima accident (40). In the Russian Federation, the Dimitrovgrad station indicates that the residence time of the plume evaluated by simulation is less than 24 h. However, the simulated hourly concentration reaches values between  $500 \text{ mBq/m}^3$  and  $1 \text{ Bq/m}^3$  on 27 September, 5 times higher than the maximum simulated in Romania (Bucharest and Constanta stations). The sampling period at the Dimitrovgrad station, which lasts for 10 d, explains the much lower levels measured there than in Romania.

At the stations in Romania, the country where the highest concentrations were measured, the level of agreement between the simulations and the observations is satisfactory even though the observed maximum concentrations are slightly underestimated in the western part of Romania.

Daily or twice-daily sampling at the Romanian stations makes it possible to realistically evaluate the residence time of the plume, which did not exceed 48 h. The simulations manage to reproduce this residence time fairly accurately. It can be seen that the simulated arrival time of the plume at the Bucharest station is several hours late, which may be related to the daily resolution of the reconstructed source term. In central Europe, the simulations show that the plume stayed for  $\sim 72 \text{ h}$  between 1 and 4 October (Krakow station in Poland). The vast majority of the simulated and observed concentrations are within a factor of 5. The simulated hourly concentrations vary between 20 and  $50 \text{ mBq/m}^3$ .

In Greece, the Athens station (Fig. 5) indicates that the simulated plume was present between 9 and 11 d. The simulated hourly concentrations, although lower in central Europe, reach  $20 \text{ mBq/m}^3$ . Finally, further west in Europe, the highest concentrations are reported in Italy and sometimes exceed  $50 \text{ mBq/m}^3$  at the Udine station during the day on 3 October. At this station, the event is relatively well reconstructed by the simulations, although the plume residence time and concentrations are slightly underestimated. At the Perugia station and at other Italian stations further west, the underestimation of the concentration levels observed becomes very large (by a factor of 3 to 10). This underestimate could be induced by the effects of the Alpine relief insufficiently accounted for by the meteorological model at a spatial resolution of  $0.5^\circ$ . It should also be added that the dispersion error increases the further one moves from the potential source of the release.

**Comparison with  $^{106}\text{Ru}$  Deposition Measurements.** A deposition measurement is generally subject to more variability than a measurement of air concentration because of wet or dry deposition patterns. Its use in models also requires realistic modeling of physical processes such as the deposition of radionuclides by wet scavenging. The deposition measurements were not used to estimate the inverted source term. As a result, comparing the simulated deposits with these observations is a relevant way of validating the source term, especially the total amount assessed. A direct analysis of the observations for deposition indicates that

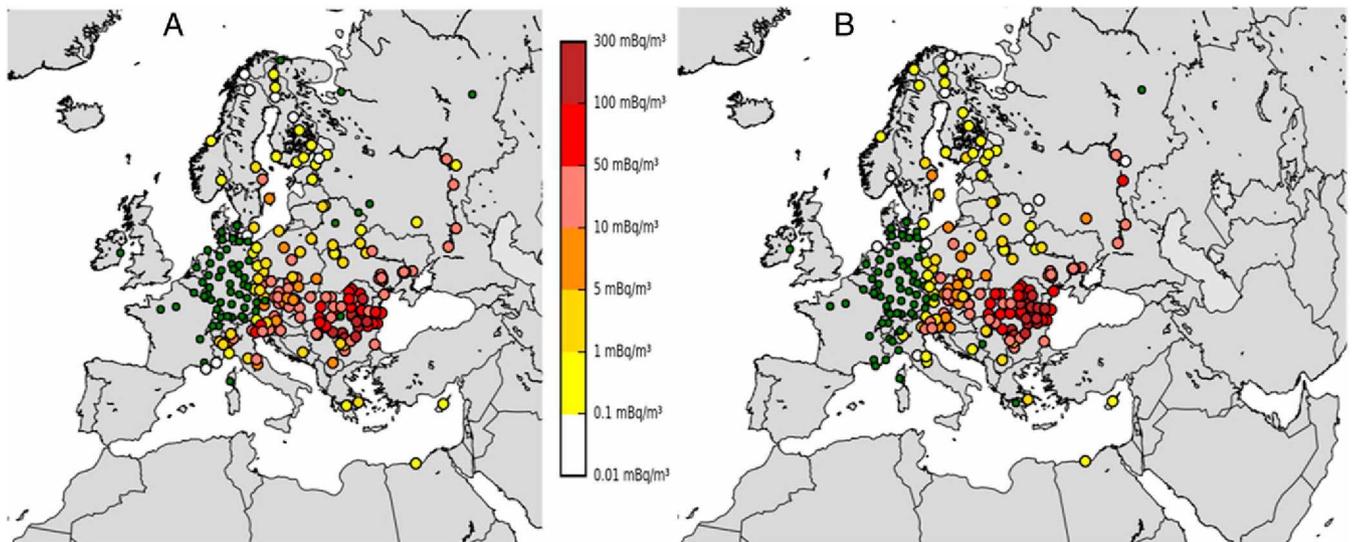


**Fig. 3.** Dispersion of the  $^{106}\text{Ru}$  plume assuming a release from the Mayak site. (A) 28 September 1200, (B) 29 September 1200, (C) 1 October 0000, (D) 2 October 1200, (E) 3 October 1200, (F) 7 October 1200. The dots represent the observed concentrations. Measurements under detection limit are represented by green dots. Precipitations are ranged from 0.2 and  $10 \text{ mm}\cdot\text{h}^{-1}$ .

southern Ural is the first region where  $^{106}\text{Ru}$  was detected on 23 September. This corroborates the results obtained by inverse modeling targeting southern Ural as the source area of the  $^{106}\text{Ru}$  and identifying a small release on 23 September. The simulation

of total deposits of  $^{106}\text{Ru}$ , based on the reference source term, is shown in Fig. 6.

According to Fig. 6, the most significant simulated deposits were concentrated in southwestern Ural, southern Scandinavia,

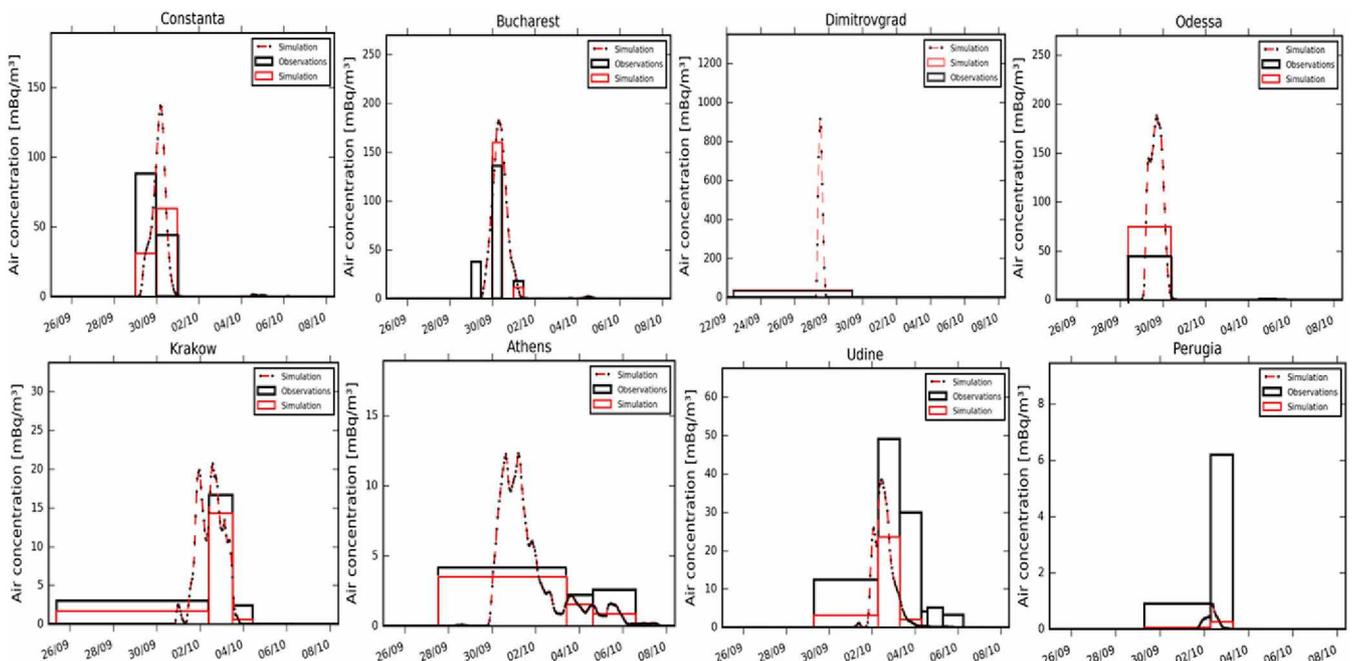


**Fig. 4.** Maximum  $^{106}\text{Ru}$  air concentrations for each monitoring station. (A) Observations. (B) Forward simulation forced using reconstructed source term from Mayak. Green dots represent air concentration measurements below the detection limit.

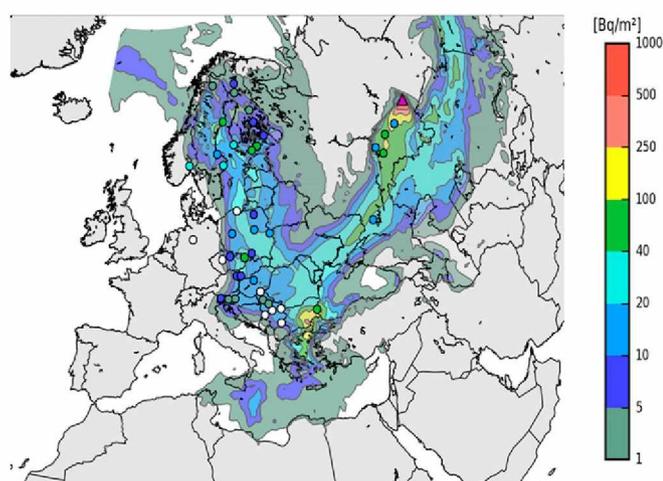
and southeastern Bulgaria, with cumulative quantities sometimes higher than  $100 \text{ Bq/m}^2$ .

Fig. 7 compares the observed surface activity of  $^{106}\text{Ru}$  with the temporal evolution of the simulated deposition. Since the IdX model is performed using  $0.5^\circ$  spatial resolution, several deposition measurements in southern Ural, located in the mesh grid containing the source location, were not exploited. Therefore, only stations located beyond a distance of 200 km from the Mayak site have been taken into account in the model-to-data comparison. Several of these stations are located in western Ural (Bugulma and Dema-Ufa stations) around 500 km west-southwest of the Mayak site. The simulations have a tendency to overestimate the observed deposits by about a factor of 3, partic-

ularly in Dema-Ufa. However, the data are incomplete at this station, since daily deposition values are available only for 26 September. Other Russian stations are situated on the edge of the most significant deposits in an area of steep gradients (Fig. 6), which explains the difficulty of reproducing the deposits observed at these stations. Low deposit levels are recorded at several Russian stations including Dimitrograd at the beginning of October, despite somewhat limited air concentrations levels. The deposition simulations have difficulty in reproducing this event, despite simulated air concentrations that coincide rather well with the observed air concentrations. The simulated deposits in southeast Bulgaria are  $> 100 \text{ Bq/m}^2$ . These simulated deposition values are sometimes greater than those simulated in western Ural. The



**Fig. 5.** Comparisons between simulated and observed air concentrations at several stations in Europe. A black rectangle shows the activity observed and the red rectangle the simulated air concentration for the observation period. The red dash represents the simulated hourly concentration.



**Fig. 6.** Map of simulated cumulative  $^{106}\text{Ru}$  deposition using the reconstructed source term from Mayak. Dots represent the observed cumulative  $^{106}\text{Ru}$  deposition. White dots represent measurements under detection limit.

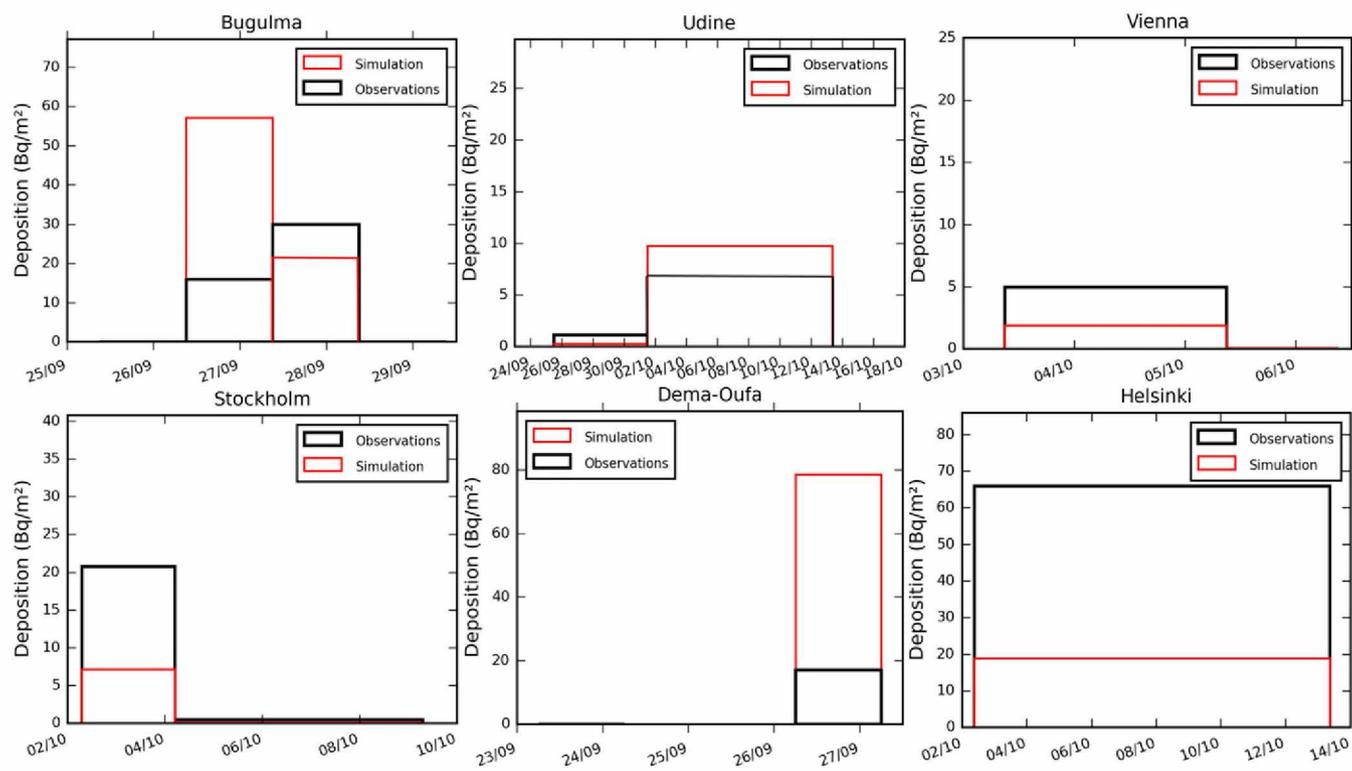
origin of these larger deposits is explained by the occurrence of precipitation in southeastern Bulgaria as the plume was passing over between 30 September and 1 October. No rainfall is modeled at the Russian stations in southwest Ural when the main plume passed over between 26 and 28 September. To confirm the pertinence of the simulated deposits in Bulgaria, plant and soil samples were taken in several villages in southeast Bulgaria (1) at the start of October 2018. All plant samples and especially litter attest of the presence of  $^{106}\text{Ru}$  deposition. However, because of the difficulty to convert becquerels per kilogram into becquerels per square meter, only the soil sample was used, and it confirmed the pres-

ence of  $^{106}\text{Ru}$  with measured deposits of between 50 and 500  $\text{Bq}/\text{m}^2$ , which is consistent with the simulations. Southern Scandinavia is characterized by measured deposits of nearly 100  $\text{Bq}/\text{m}^2$ , particularly in southern Finland (Helsinki station). These deposits seem to be due to heavy rainfall as the plume was passing over during daytime on 2 October. Although the observed air concentrations are reproduced satisfactorily throughout Scandinavia by the simulations, the observed deposits are underestimated by the simulations, especially in southern Finland.

In western Europe, the agreement between simulated and observed deposits is more satisfactory (Udine station, Italy), despite the air concentrations' being underestimated. Finally, low deposition values ( $<5 \text{ Bq}/\text{m}^2$ ) were reported in the Czech Republic and Serbia; these deposits were overestimated by the simulations. The scores are not as good as for air concentration since the proportion of the simulated deposition that is within a factor of 5 of the observed values is 52%.

## Conclusion

In this paper, we used an inverse modeling method to identify the origin of  $^{106}\text{Ru}$  detections measured in Europe in fall 2017 and to estimate the total amount released into the atmosphere. The method is part of a variational approach, which remains the most appropriate one for operational use. More than 1,100 air concentration observations were used in the inverse modeling process. The observation errors were modeled using log-normal statistics, whereas the background errors were assumed to be Gaussian by enforcing source positivity. The obtained results indicate that a release of about several hundred terabecquerels, emitted from the southern of the Urals region, is the most likely hypothesis. It is a release from the southern Urals that manages to consistently explain all of the observations. The meteorological conditions, mainly dominated by the presence of an anticyclone over northeastern Scandinavia, encouraged the dispersion



**Fig. 7.** Comparisons between simulated and observed deposits at several stations in Europe. A black rectangle shows the cumulated observed activity over a period and a red rectangle the simulated deposition for the observation period.

of the  $^{106}\text{Ru}$  from southern of the Urals toward western Europe between the end of September and the very beginning of October. Of the known nuclear facilities in this region that are likely to release large quantities of  $^{106}\text{Ru}$ , the Mayak site is the most pertinent. Forward simulation of a release from Mayak consistently reproduces observations of air concentrations in Europe and it is the only identified site compatible with the deposition observations made in the southern Urals. The source term evaluated from the Mayak site is about 250 TBq. The density of the observations means that a regularization term is not necessary when resolving the inverse problem, guaranteeing a robust estimate. The simulations made using the source term from Mayak indicate that the percentage of observed and simulated concentrations within a factor of 2 exceeds 40%. The percentage of observed and simulated observations within a factor of 5 is greater than 70%. It can be seen, however, that the observations made furthest from the release point, in Italy, are the most difficult for the simulations to reproduce. Comparisons of the simulations with the deposition measurements, not taken into account in the inversion process, are also realistic. The simulations identify the occurrence of several rainfall events, which are the cause of the

higher  $^{106}\text{Ru}$  deposits measured in Bulgaria and Scandinavia. Several samples taken in Bulgaria confirmed the presence of  $^{106}\text{Ru}$  in proportions similar to those calculated by simulation. However, the deposit is sometimes underestimated by the simulations, particularly in southern Scandinavia. The majority of the release would have occurred during the day on 26 September. Other smaller releases could also have occurred from 23 September, consistent with positive deposition measurements from 23 September in the southern Urals. In future, the use of meteorological fields with a finer spatial and temporal resolution is envisaged, so that observations within a 50-km radius of the release location can be used. Integrating a more realistic wet deposition model (41) could also improve the reconstruction of the  $^{106}\text{Ru}$  deposition observed in southern Scandinavia. Finally, IRSN is currently working on the development of probabilistic MCMC methods (42). This type of approach complements the variational approach because it enables the uncertainties of the result of inverse modeling obtained using a variational approach to be quantified.

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